

四川盆地中部合川—潼南地区茅二下亚段白云岩储层特征及成储机制

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摘要 近年来四川盆地中部合川—潼南地区中二叠统茅二下亚段白云岩储层勘探取得了重要突破,但储层强非均质性制约了有利区带的预测。基于岩心和薄片观察、物性实验及地球化学分析,系统剖析茅二下亚段白云岩储层的基本特征,讨论白云岩的成因和储层形成机制。结果表明:①茅二下亚段储集岩以针孔状白云岩与孔洞状白云岩为主,储集空间类型为晶间(溶)孔、溶洞和裂缝;储层类型为以裂缝-孔洞型为主,裂缝-孔隙型为辅;以片状喉道为主,孔喉配置关系中等,裂缝为主要的渗流通道;针孔状白云岩平均孔隙度为2.50%,平均渗透率为 $0.806 \times 10^{-3} \mu\text{m}^2$,孔洞状白云岩平均孔隙度为4.32%,平均渗透率为 $1.047 \times 10^{-3} \mu\text{m}^2$ 。②高能滩相是白云岩储层形成的物质基础,多期岩溶作用是形成优质储层的关键。准同生期白云石化作用增强了岩石抗压实能力,有助于储层孔隙的保持;茅口组沉积末期热液白云石化以胶结充填先期孔隙为主,降低了白云岩的储集性能。

关键词 白云岩; 储层特征; 成储机制; 茅口组; 四川盆地

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0 前言

中二叠统茅口组是四川盆地增储上产的重要层系。前期勘探主要针对茅口组顶部与风化壳相关的灰岩缝洞型储层,在川南地区发现了一批与现今构造相关的油气藏群^[1-3]。随着南充1井、角探1井等高产井的突破^[4],茅二上亚段白云岩成为四川盆地茅口组勘探的重点。该套白云岩受15号断裂带控制,具有北西—南东向展布的特征,向北可达梓潼地区,向南可达卧龙河—丰都地区^[5-6]。针对茅二上亚段白云岩储层成因,学者们基于有限的钻井资料,提出了构造-热液^[7-10]、热次盆^[6, 11]等白云石化作用与东吴运动末期暴露溶蚀作用^[4, 12]、埋藏溶蚀作用^[13]共同控制储层发育的观点。近年来,中国石油大庆油田有限责任公司在合川—潼南区块钻遇

茅二下亚段白云岩储层,多口井获高产工业气流^[14-15],并且试采效果较好,预示着该套白云岩储层具有巨大的勘探开发潜力。然而在后期评价井实施阶段发现,该套白云岩储层非均质性较强,横向变化规律不清,亟需深入认识储层成因机制。

本文基于合川—潼南地区茅二下亚段最新的连续取心井资料,开展储层岩石学、地球化学工作,厘清白云石化与溶蚀作用之间的关系,并进一步分析其成储机制,建立了滩控岩溶型白云岩储层发育模式,期望为加快茅二下亚段白云岩储层的勘探开发提供理论支撑。

1 茅口组沉积特征

在二叠纪栖霞末期海平面下降之后,茅口早期海平面开始缓慢上升,形成了一套颜色相对较深的

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眼球眼皮状灰岩(纹层状泥灰岩与生物碎屑泥晶灰岩互层)为主的台地相沉积,其后进入成滩期(茅二段—茅三段沉积期)。茅二期四川盆地以半局限台地环境为主,台内滩广泛发育,在盆地西北缘开始出现台地边缘沉积^[16](图1a)。茅三期整体海平面较为开阔,形成了一套以生物碎屑灰岩为主的高能

滩相沉积^[14]。

研究区主体位于川中缓褶皱带的合川—潼南地区,面积约为4 000 km²(图1b)。区内茅二下亚段主要发育相对高能的台内滩相白云岩、白云质灰岩以及生物碎屑灰岩(图1c);茅二上亚段以泥晶生物碎屑灰岩为主,代表海平面相对上升的沉积。

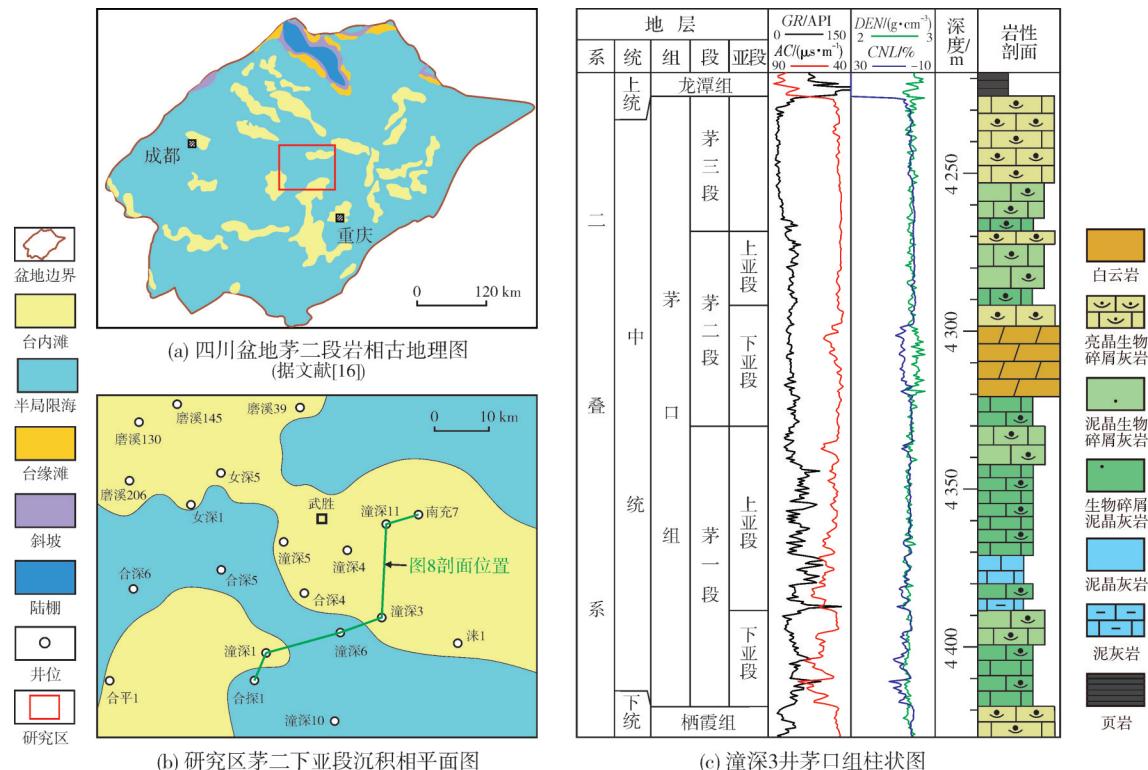


图1 四川盆地合川—潼南地区茅二段沉积背景及地层综合柱状图

Fig.1 Sedimentary background of Maokou Member 2 and stratigraphic column of Maokou Formation in Hechuan-Tongnan area, Sichuan Basin

2 储层特征

2.1 储集岩类型

研究区茅二下亚段白云岩分为致密细晶白云岩(图2a)、针孔状白云岩(图2b, 2c)与孔洞状白云岩(图2d—2f)。其中,后二者是研究区茅二下亚段最主要的储集岩类。针孔状白云岩岩心呈黑灰色,断面粗糙,可见大量非均匀分布的针孔(图2b);显微镜下观察,矿物主要以自形—半自形形状细—中晶白云石为主(图2c),白云石具雾心亮边结构。孔洞状白云岩岩心呈黑灰色,发育大量残余溶蚀洞穴,溶洞直径大小不一(图2d, 2e);矿物类型主要为半自形形状中—细晶白云石,可见雾心亮边结构,发育颗粒幻影,靠近孔洞边缘白云石晶粒自形程度更高,晶粒更粗(图2f)。

2.2 储集空间类型

对合川—潼南地区茅二下亚段岩心及薄片进行观察和统计,发现茅二下亚段白云岩类储集体的储集空间类型丰富(图3),以晶间(溶)孔、溶洞为主,可见裂缝。

晶间(溶)孔 晶间(溶)孔直径在0.1~1 mm之间,常发育于针孔状白云岩中,岩心上由于溶蚀作用呈现出针孔状(图3a)。显微镜下可见晶间孔由自形程度相对较好的白云石晶粒边缘围限而成(图3b);若存在溶蚀作用,其边缘棱角会逐渐模糊,边界呈港湾状,形成晶间溶孔。镜下多见颗粒幻影,孔隙发育于颗粒幻影之间,推测其原岩为相对高能的颗粒灰岩,而孔隙空间则基本继承于原始粒间孔隙。

溶洞 为研究区重要的储集空间类型之一,多

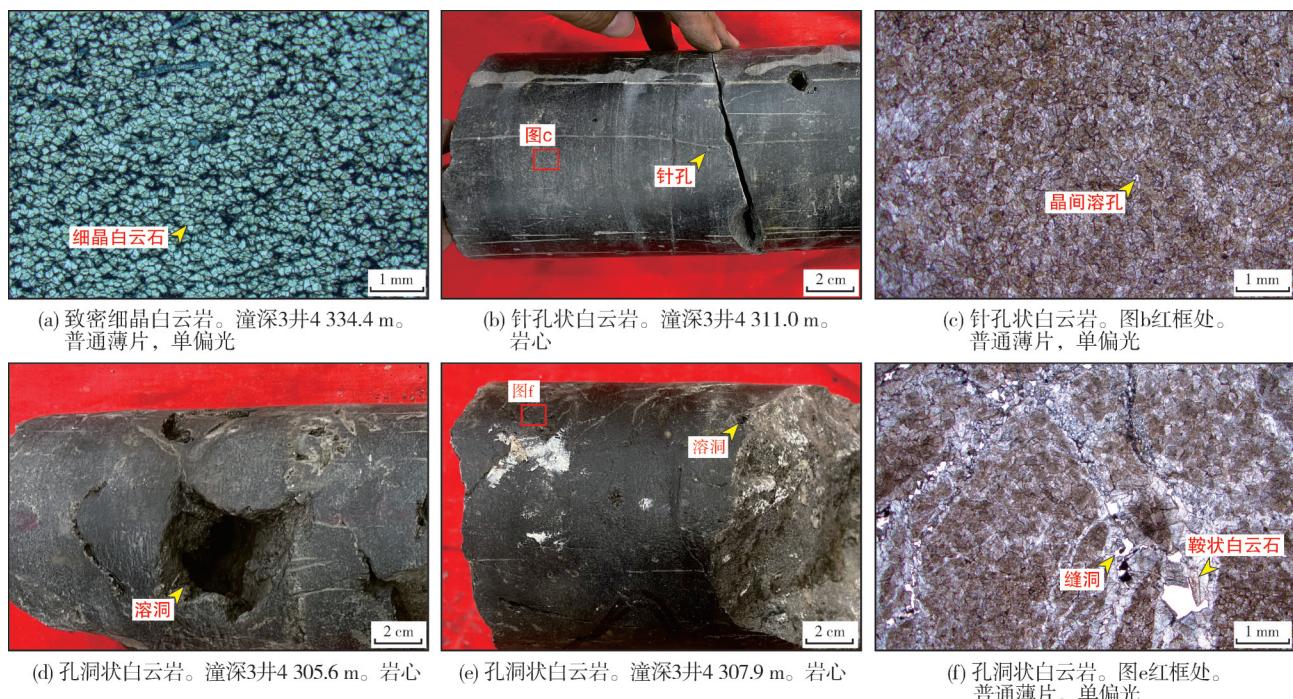


图2 合川—潼南地区茅二下亚段白云岩岩石学特征

Fig. 2 Petrological characteristics of dolomite of the lower submember of Maokou Member 2 in Hechuan-Tongnan area

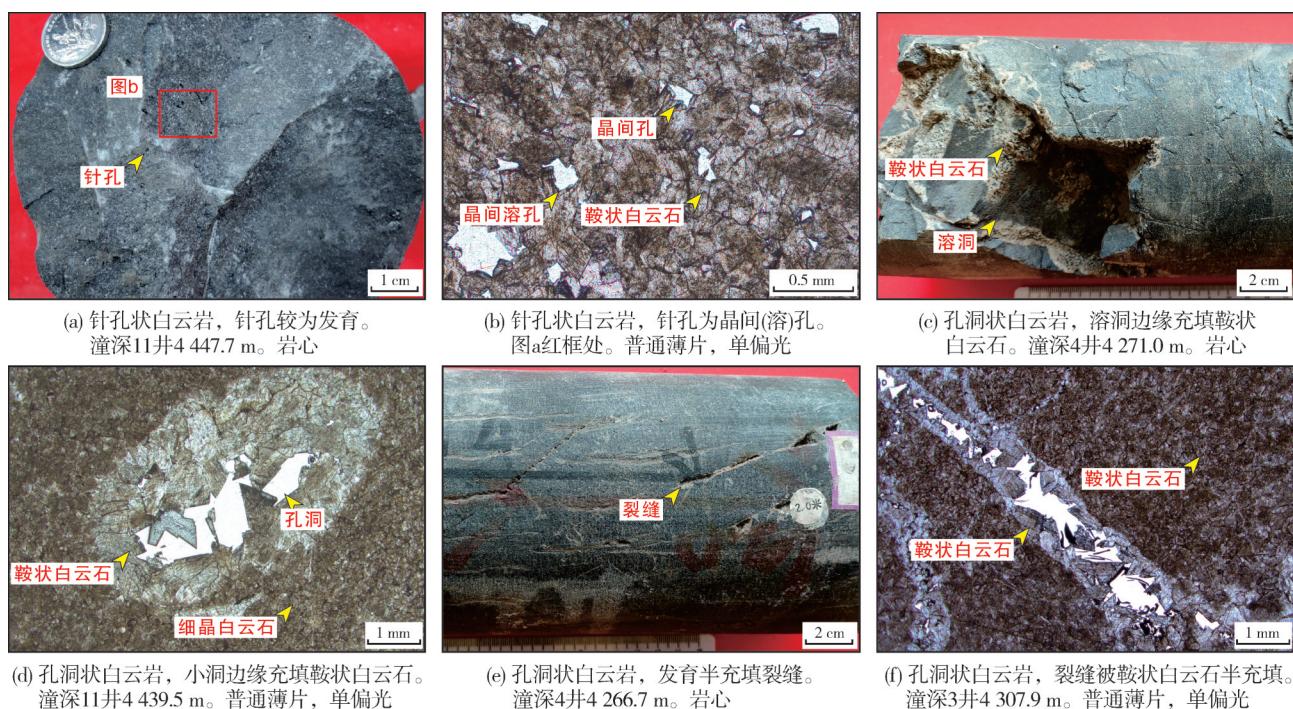


图3 合川—潼南地区茅二下亚段白云岩储集空间特征

Fig. 3 Reservoir space characteristics of dolomite of the lower submember of Maokou Member 2 in Hechuan-Tongnan area

见于孔洞状白云岩中(图3c),溶洞尺寸大小不一,其直径在0.2~10 cm之间。溶洞边缘常发育有巨粗晶鞍状白云石胶结物,将溶洞半充填或全充填(图3c,

3d)。基岩中的中—细晶白云石多见溶蚀特征,而鞍状白云石的溶蚀特征不明显(图3d)。溶洞多为早期原生孔隙或裂缝经持续性扩溶而形成。

裂缝 包括构造缝和压溶缝2种类型。构造缝一般为高角度缝(图3e),低角度缝相对较少。构造缝边界相对平直,部分存在明显的溶蚀特征,可与溶洞交织出现;多被粗晶鞍状白云石半充填或全充填(图3f),缝隙之间有时可见少量沥青。由于白云岩具有较好的抗压溶压实的特征,其内部压溶缝相对较少,多出现于不同岩性的交界处。

2.3 物性特征

对研究区3口井(潼深3、潼深11、潼深4)的156

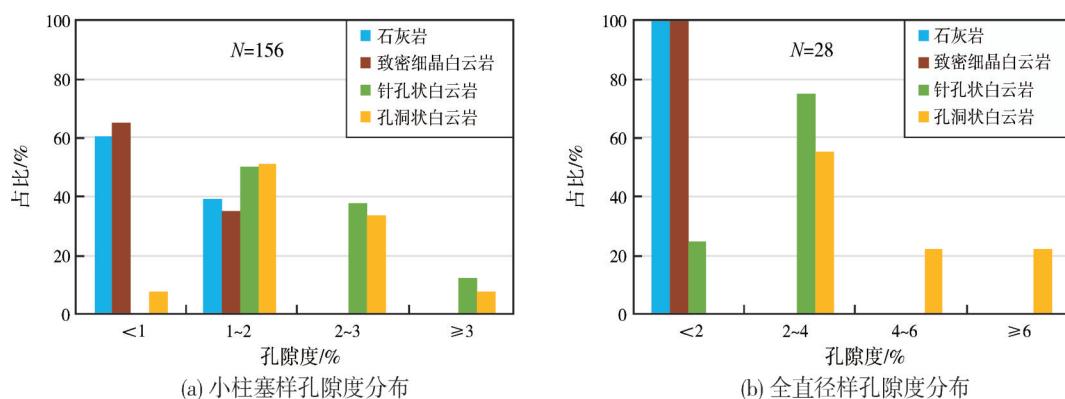


图4 合川—潼南地区茅二下亚段碳酸盐岩孔隙度分布直方图

Fig. 4 Porosity histograms of carbonate rocks of the lower submember of Maokou Member 2 in Hechuan-Tongnan area

为了精细分析不同类型白云岩储层参数的差异,选取以上样品分岩性绘制茅二下亚段孔渗交会图(图5)。由图5可知:石灰岩与致密细晶白云岩的孔隙度普遍低于2%,致密细晶白云岩中渗透率大于 $1 \times 10^{-3} \mu\text{m}^2$ 的占比略高于石灰岩;作为最主要的

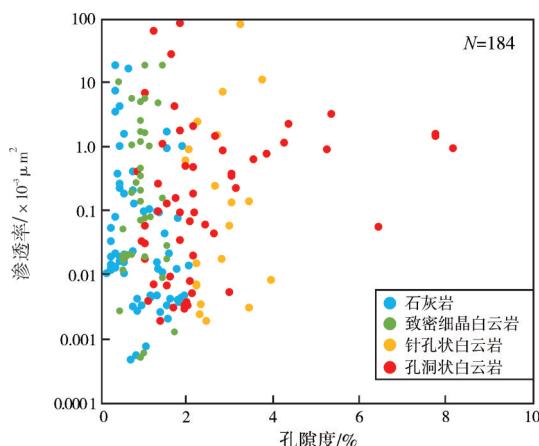
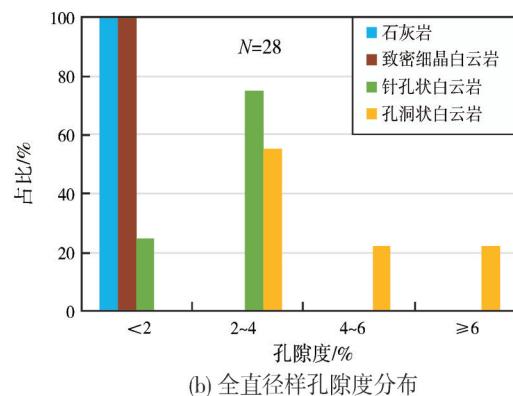


图5 合川—潼南地区茅二下亚段碳酸盐岩孔隙度-渗透率交会图

Fig. 5 Porosity-permeability crossplot of carbonate rocks of the lower submember of Maokou Member 2 in Hechuan-Tongnan area

个柱塞样的物性与28个全直径样的物性进行了统计分析。柱塞样品中:致密细晶白云岩孔隙度主要分布在0~2%之间,而针孔状白云岩和孔洞状白云岩的孔隙度主要分布在1%~3%之间(图4a)。全直径样中:石灰岩、致密细晶白云岩与孔洞状白云岩、针孔状白云岩的孔隙度差异更明显(图4b),孔洞状白云岩的孔隙度分布特点反映更充分;针孔状白云岩平均孔隙度为2.50%,平均渗透率为 $0.806 \times 10^{-3} \mu\text{m}^2$;孔洞状白云岩呈现出高孔隙度特征,孔隙度最大值为8.11%,平均值为4.32%,平均渗透率为 $1.047 \times 10^{-3} \mu\text{m}^2$ 。



储集岩,针孔状白云岩与孔洞状白云岩的孔隙度普遍大于2%,占比达到58.45%;孔洞状白云岩与针孔状白云岩大部分样品孔渗具有良好的正相关性,前者比后者具有更优的孔隙度与渗透率。在图5中,一部分样品在孔隙度变化不大的情况下渗透率急剧增加,表明裂缝对区内白云岩储层的影响较为显著。结合白云岩储层宏观特征,认为研究区茅二下亚段白云岩储层主要以裂缝-孔洞型为主、裂缝-孔隙型为辅。

2.4 孔隙结构特征

茅二下亚段白云岩储层以片状喉道为主,根据孔喉参数特征可划分为2种主要类型(图6):

凹形压汞曲线 主要出现在针孔状白云岩中,表现为曲线中间的过渡段较平缓,汞不饱和体积常小于10%,退汞量约为30%,具有一定的粗歪度,孔喉半径中值为 $0.185 \mu\text{m}$,饱和中值压力为4.07 MPa。此类压汞曲线反映了储层的孔隙较为发育,孔径分布较为均一,孔隙类型较为单一,孔喉连通性中等(图6a)。

台阶形压汞曲线 主要出现在孔洞状白云岩

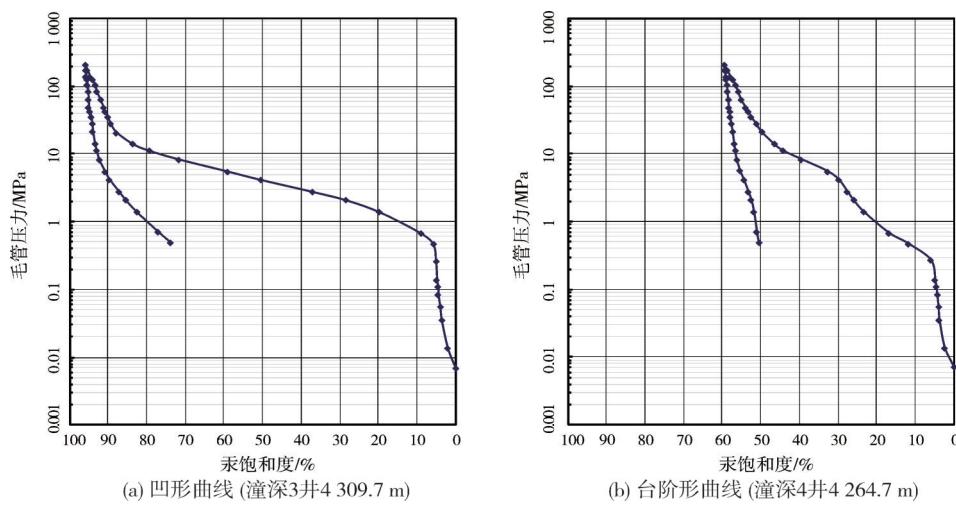


图6 合川—潼南地区茅二下亚段白云岩储层典型压汞曲线特征

Fig. 6 Typical mercury intrusion curves of dolomite reservoirs of the lower submember of Maokou Member 2 in Hechuan-Tongnan area

中,毛管压力曲线形态陡峭,初始段、中间平缓段和末端上翘段无明显区分界限,中间平缓段短,汞不饱和体积常大于40%,退汞量约为10%,具有一定的细度,孔喉半径中值为0.034 μm,饱和中值压力为21.77 MPa。此类压汞曲线反映了储层的孔隙较少,孔径分布不均,孔隙、孔洞均发育,孔喉连通性较差(图6b)。

3 白云岩储层主控因素

3.1 颗粒滩是储层形成的物质基础

典型取心段垂向发育序列分析表明:颗粒滩单旋回自下而上往往发育生物碎屑泥晶灰岩-泥晶生物碎屑灰岩-白云岩的岩性组合,白云岩储层通常分布于颗粒滩旋回的中上部(图7),主要为针孔状

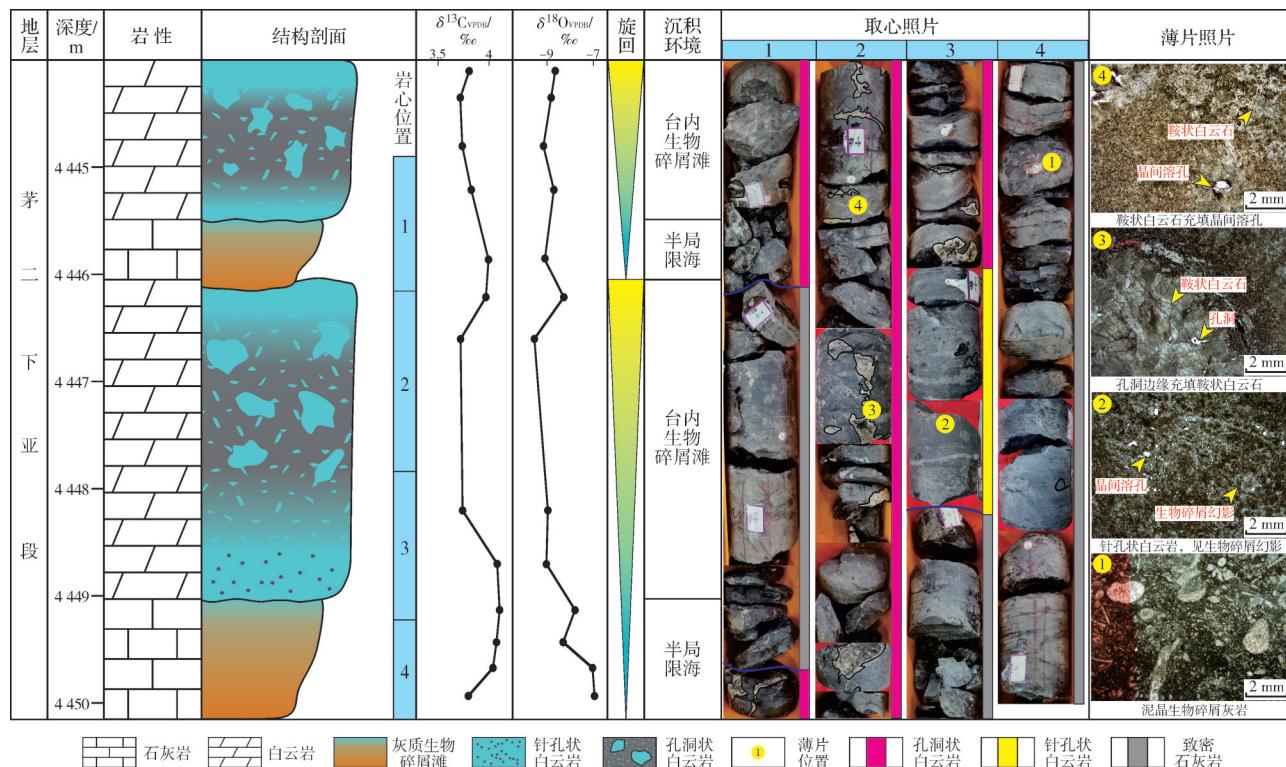


图7 合川—潼南地区茅二下亚段白云岩垂向发育序列(潼深11井)

Fig. 7 Vertical dolomite sequence of the lower submember of Maokou Member 2 in Hechuan-Tongnan area (Well Tongshen 11)

白云岩、孔洞状白云岩。由于颗粒滩沉积旋回顶部受到准同生期暴露岩溶作用的影响, 碳、氧同位素组成具有自下而上逐渐负偏移的特征^[17-18](图7), 到下一旋回初期又出现明显正偏移, 这反映了典型的向上变浅的浅水颗粒滩沉积环境。通过研究区茅二下亚段沉积相连井对比分析发现, 白云岩储层主要发育在台内滩亚相中, 如潼深3井至南充7井一带台内滩规模较大, 白云岩储层分布较连续(图8)。由此可见, 沉积环境控制了储层的发育。

研究区茅二下亚段白云岩储层在镜下常见有明显的颗粒幻影, 表明其原岩为相对高能的颗粒灰岩。首先, 颗粒滩在沉积期具有良好的原始孔隙以及大量不稳定矿物溶蚀孔, 生物碎屑颗粒常呈点-

线接触, 而且准同生期胶结程度低, 可为白云石化流体提供有利运移通道, 从而促使白云石化作用在颗粒滩沉积体中持续进行^[19-20]。其次, 茅口组沉积期气候变化剧烈, 具有多幕式冰期存在^[21-22], 强烈的气候变化使得海平面波动剧烈, 颗粒滩常常由于可容空间不足而叠置迁移(图8), 从而导致局部水体受限, 促使水体盐度变高, 为早期白云石化提供了一定的流体环境。再者, 颗粒滩往往发育在水体较浅的地貌高部位, 向上变浅旋回的中上部容易遭受淡水溶蚀作用, 从而发育一系列溶蚀孔洞^[23-24], 改善滩体的储渗条件。综上所述, 颗粒滩为茅二下亚段白云岩储层形成提供了重要的物质基础。

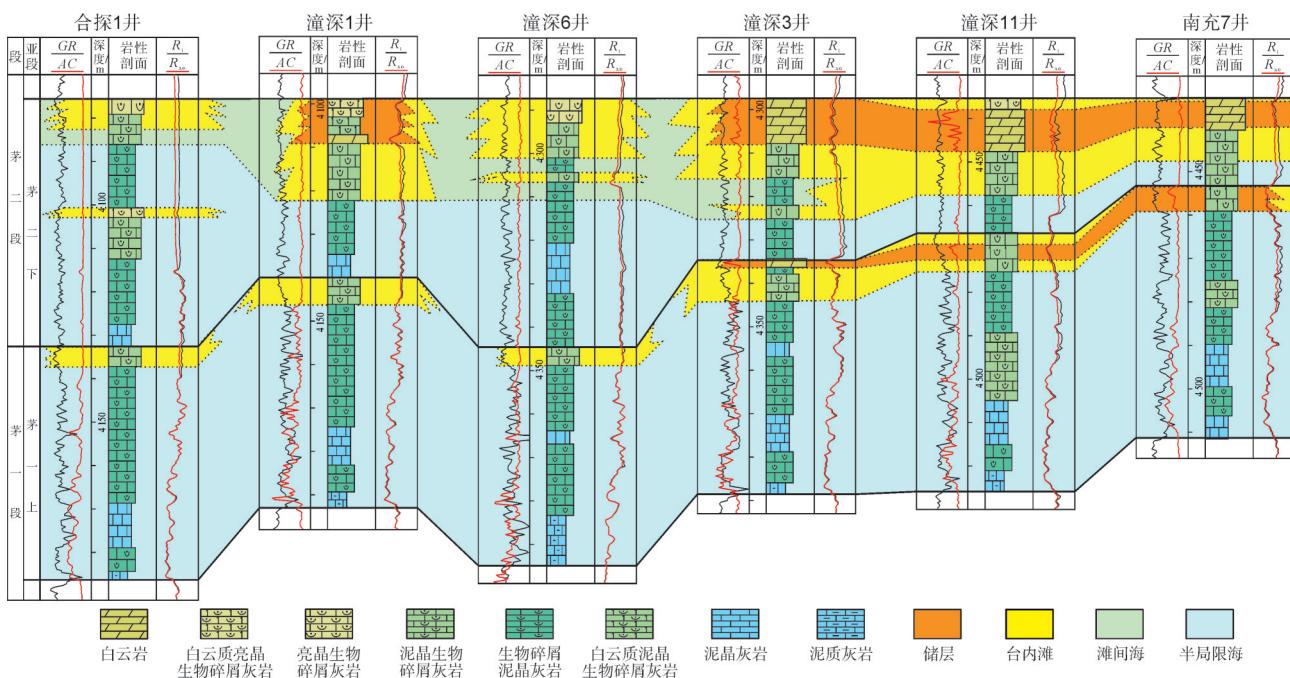


图8 合川—潼南地区茅口组沉积相与白云岩储层连井对比图(剖面位置见图1b)

Fig. 8 Inter-well comparison section of sedimentary facies and dolomite reservoir of Maokou Formation in Hechuan-Tongnan area
(section location is shown in Fig. 1b)

3.2 多期岩溶作用对储层的差异改造

颗粒滩不仅是白云岩储层重要的物质基础, 也是早成岩期岩溶作用发生的有利场所^[23]。早成岩阶段岩石往往具有较好的孔渗性^[25], 尤其在颗粒滩环境中岩石自身具有更好的孔渗条件, 可以为岩溶水提供流通通道。四川盆地茅口组滩相沉积中广泛发育各种类型的岩溶系统, 是典型的早成岩期岩溶作用的产物^[26]。其中, 准同生期岩溶作用是早成岩早期发生频率更高的岩溶作用, 其发育主要受到

高频旋回的控制, 旋回中上部岩溶作用往往较发育, 而旋回下部岩溶作用相对较弱^[18,27]。

研究区茅二下亚段发育大量高频向上变浅序列, 滩相白云岩存在明显的岩溶差异改造作用。合川—潼南地区茅口组不同取心井或者同一钻井不同地段的岩溶作用强度均具有较大差异(图9):①当准同生岩溶作用较强时, 旋回上部受到岩溶系统切割而角砾化; 旋回中部由于潜流溶蚀的原因形成大量规模不等的溶洞, 溶洞边缘可见鞍状白云石胶结充填; 旋回下部偶见小型溶洞和针孔。优质储层主要发

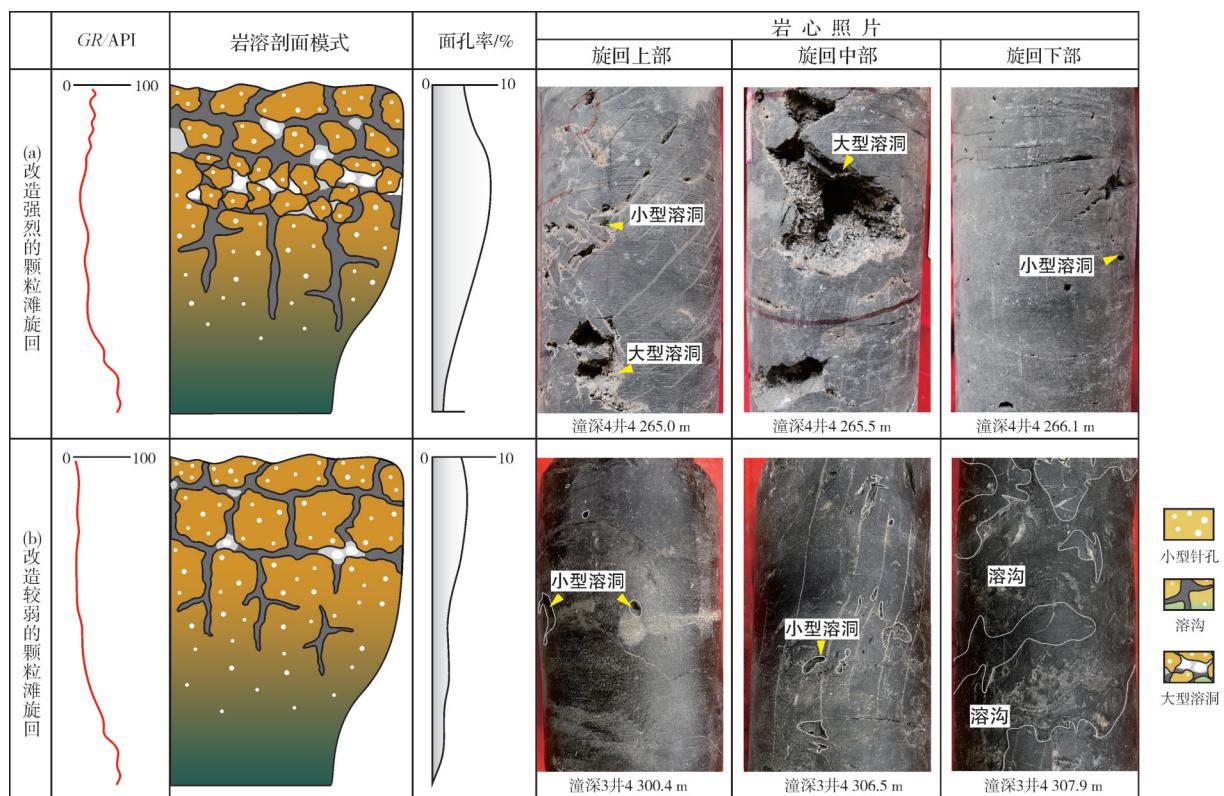


图9 颗粒滩暴露岩溶作用与储层的关系

Fig. 9 Relationship between the high frequency exposure karstification and reservoirs

育于旋回的中部,上部由于岩溶过于强烈而导致储层质量降低。②当准同生岩溶作用较弱时,旋回中上部发育少量溶沟,并被碳酸盐岩砂与不溶残余混合充填,岩石整体以基岩为主,并发育针孔以及扩溶形成的小型溶洞。优质储层主要发育于旋回的中上部,越接近暴露面,储层质量越好。

茅口组沉积中期,冰期作用导致大规模海平面下降^[21-22,28](已在川西北地区得到证实^[29]),研究区茅二下亚段遭受大规模的早成岩期淡水溶蚀作用。合川—潼南地区茅二下亚段白云岩在准同生期岩溶作用与早成岩期岩溶作用的共同影响下,形成以孔洞为主、孔隙为辅的规模性的白云岩岩溶储层。

3.3 多期白云石化作用对储层的保持与破坏

研究区茅二下亚段石灰岩和白云岩碳氧同位素组成总体落在二叠纪海水范围,并且致密细晶白云岩 $\delta^{18}\text{O}$ 值较石灰岩明显偏正,表明其受到高盐度流体的影响,而岩溶改造后的白云岩的 $\delta^{13}\text{C}$ 、 $\delta^{18}\text{O}$ 值较石灰岩则相对偏负。 $\delta^{13}\text{C}$ 负偏可能是由于大气淡水影响导致,而 $\delta^{18}\text{O}$ 负偏则可能是受热液活动或淡水影响(图10a)。将区内碳酸盐岩稀土元素采用

PAAS标准化后发现,白云岩类的稀土元素配分形态近似,均为“轻稀土亏损、重稀土富集”的左倾形态(图10b),与同期石灰岩的稀土元素配分形态类似,而且有别于鞍状白云石的Eu正异常,表明白云石化流体主要为海源流体。一些学者通过白云岩U-Pb同位素测年法得到茅口组白云岩形成的时间为 $257\text{Ma} \pm 10\text{ Ma}$ ^[14,30],大致确定了茅口组白云石化作用发生的时间为二叠纪。上述结果反映研究区茅二下亚段白云石化流体性质与沉积期海水性质类似,并且为高盐度流体,区内白云岩整体为准同生期渗透回流白云石化成因。

尽管越来越多的证据表明中二叠统白云岩的形成时间主要在准同生期,但中二叠世末期东吴运动引发的构造-热液白云石化作用仍不可否认,这一期白云石化作用加剧了白云岩储层的非均质性。中二叠世末期峨眉地裂运动进入高潮期^[28,31],大量富镁离子的幔源热液顺着深大断裂涌入茅口组,主要表现为热液白云石化流体以白云岩早期溶洞、裂缝为优势通道,在孔洞边缘形成鞍状白云石胶结物,从而充填白云岩储层的先期孔隙,对储层起到破坏作用。孔隙、孔洞以及裂缝边缘的鞍状白云石,具

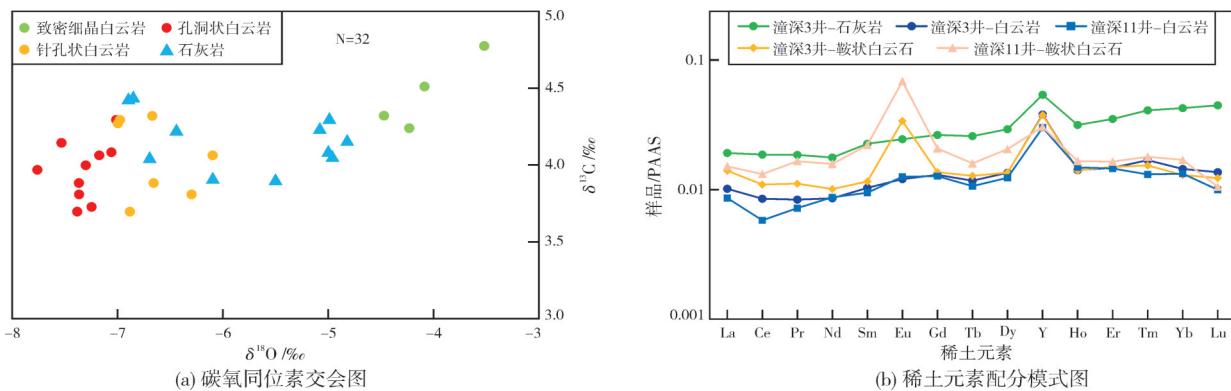


图 10 合川—潼南地区茅二下亚段碳氧同位素分布与稀土元素配分形态图

Fig. 10 The $\delta^{18}\text{O}-\delta^{13}\text{C}$ crossplot and rare earth element distribution patterns of the lower submember of Maokou Member 2 in Hechuan-Tongnan area

有与基质白云石完全不同的地球化学特征(图10b),反映白云石化流体与幔源热液密切相关。合川—潼南地区茅口组白云岩孔隙主要为继承于原始颗粒灰岩的粒间孔隙,早期白云石化提供了更加抗压溶压实的岩石骨架,对储层的作用是以保持孔隙为主,而茅口组沉积末期的热液白云石化则主要是破坏性作用,这与栖霞组白云石化作用有相似之处^[32]。

4 白云岩储层发育模式

综上所述,四川盆地合川—潼南地区茅口组主要发育滩控岩溶型白云岩储层,总体上储层发育主要受沉积期地貌高地颗粒滩相、多期岩溶作用以及2期白云石化作用的共同控制。这三者对白云岩储层的形成具有相互影响、交织发育的内在联系,为此建立研究区茅口组白云岩储层发育模式,进一步厘清不同控制因素之间的相互关系。

茅口组沉积初期四川盆地以持续海侵为主,沉积物能量总体较低,随着海平面降低,茅二下亚段沉积期开始在地貌高部位发育相对高能的颗粒滩沉积(图11a)。颗粒滩岩性主要为亮晶生物碎屑灰岩、泥晶生物碎屑灰岩,颗粒类型多以生物碎屑为主,总体为中—高能沉积产物。沉积期颗粒滩沉积物初始孔隙发育,生物碎屑颗粒以点—线接触为主,普遍发育粒间孔隙。海平面下降导致滩体叠置迁移,使得滩间局限区域海水盐度增加,为早期白云石化提供流体来源(图11b)。随着海平面进一步下降,滩体暴露遭受准同生期淡水岩溶作用(图11c),岩溶水在高孔渗性的白云石化颗粒滩中漫流溶蚀,并受重力作用影响发育穿层状垂直溶沟或溶蚀斑块,在水平潜流带可发育大型溶洞。早期白云石化作用时间与准同生期岩溶作用时间相互叠合,可见

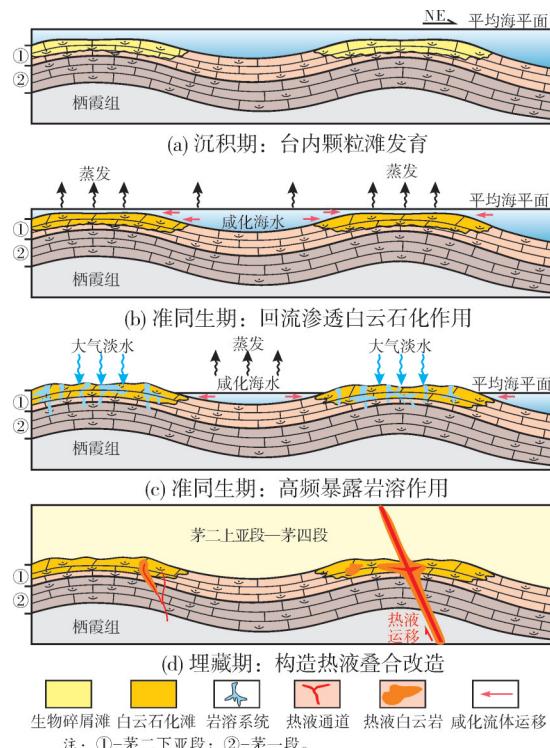


图 11 合川—潼南地区茅口组白云岩储层发育模式

Fig. 11 Development pattern of dolomite reservoir of Maokou Formation in Hechuan-Tongnan area

大量交织发育的关系。中二叠世末期东吴运动导致四川盆地大规模喷发峨眉山玄武岩^[33-34],而且研究区东北部受15号基底断裂活化影响^[5-6],大规模的热活动造成研究区内大量幔源流体顺着断裂、裂缝等通道进入茅二下亚段白云岩(图11d),在白云岩的早期孔洞中胶结充填形成鞍状白云石。远离基底断裂的白云岩受热液叠合改造影响较少,孔洞相对保存较好。

5 结 论

(1)四川盆地合川—潼南地区中二叠统茅二下

亚段发育针孔状白云岩和孔洞状白云岩2类储层,储集空间以晶间(溶)孔、溶洞与裂缝为主。全直径样分析表明:针孔状白云岩平均孔隙度为2.50%,平均渗透率为 $0.806 \times 10^{-3} \mu\text{m}^2$;孔洞状白云岩平均孔隙度为4.32%,平均渗透率为 $1.047 \times 10^{-3} \mu\text{m}^2$ 。压汞分析表明,2类储层均以片状喉道为主,整体孔隙结构中等,裂缝是主要的渗流通道。

(2)茅二下亚段白云岩储层的形成主要受到3个方面控制:①台内颗粒滩相具有良好的原始孔隙与流体形成环境,是白云岩储层形成的物质基础;②高频旋回控制的准同生期岩溶作用与茅二下亚段沉积末期的早成岩期岩溶作用改善了储集条件,形成大量溶蚀孔洞,是形成优质储层的关键因素;③准同生期白云石化作用增强了岩石抗压实能力,是储层孔隙保持的关键。茅口组沉积末期热液白云石化作用以胶结充填先期孔隙为主,降低了白云岩的储集性能。

(3)在厘清白云岩储层不同控制因素发育时序的基础上,建立了区内茅二下亚段白云岩储层发育模式,台内古地貌高地滩相发育区是此类白云岩储层的有利分布区。研究成果可为下一步盆地级的油气勘探提供指导。

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Characteristics and formation mechanism of dolomite reservoirs in the lower submember of Maokou Member 2 in Hechuan-Tongnan area, central Sichuan Basin

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Abstract: The Middle Permian Maokou Formation is an important stratigraphic series for increasing reserves and production in Sichuan Basin. The preliminary exploration mainly focused on the limestone fissure type reservoirs related to the weathered crust at the top of the Maokou Formation. In recent years, with the breakthrough of a series of high-yield wells in the central Sichuan region, the significant exploration potential in dolomite of the Maokou Member 2 has been demonstrated, including the intraplatform shoal dolomite reservoir of the lower member of Permian Maokou Member 2 in Hechuan-Tongnan area, while the strong reservoir heterogeneity restricts the further exploration. Based on core and thin section observation, physical property experiments and geochemical analysis, the basic characteristics of dolomite reservoir of the lower submember of Maokou Member 2 in Hechuan-Tongnan area are systematically analyzed, and the genesis and formation mechanism of dolomite reservoir are discussed. The results show that the dolomite reservoir of lower submember of Maokou Member 2 consists of pinhole dolomite and vuggy dolomite, and the reservoir space is characterized by intercrystal (dissolved) pore, karst cave and fracture as the main seepage channel. The reservoir type is mainly fracture-vuggy reservoir, and secondly fracture-pore reservoir. The average porosity and permeability of pinhole dolomite of the full diameter sample are 2.50% and $0.806 \times 10^{-3} \mu\text{m}^2$, and these of vuggy dolomite are 4.32% and $1.047 \times 10^{-3} \mu\text{m}^2$. Combined with previous sedimentation and diagenesis research, it is considered that the high energy shoal facies is material base for reservoir development and multi-stage karstification is the key for the formation of high quality reservoir. Penecontemporaneous dolomitization is the key to reservoir space preservation by enhancing the compaction resistance of rocks. The late hydrothermal dolomitization of Maokou Formation related with the Dongwu Movement (Emei taphrogenesis in the upper-Yangtze region) in the late Middle Permian is dominated by cementing and filling early pores, which reduces the reservoir performance of dolomite. This study provides a basis for predicting the distribution zone of high-quality reservoirs in Maokou Formation.

Key words: dolomite; reservoir characteristics; formation mechanism; Maokou Formation; Sichuan Basin

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